

# Modeling and Simulating Depositional Sequences Using Latent Gaussian Random Fields

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**Abstract** Simulating a depositional (or stratigraphic) sequence conditionally on borehole data is a long-standing problem in hydrogeology and in petroleum geostatistics. This paper presents a new rule-based approach for simulating depositional sequences of surfaces conditionally on lithofacies thickness data. The thickness of each layer is modeled by a transformed latent Gaussian random field allowing for null thickness thanks to a truncation process. Layers are sequentially stacked above each other following the regional stratigraphic sequence. By choosing adequately the variograms of these random fields, the simulated surfaces separating two layers can be continuous and smooth. Borehole information is often incomplete in the sense that it does not provide direct information about the exact layer that some observed thickness belongs to. The latent Gaussian model proposed in this paper offers a natural solution to this problem by means of a Bayesian setting with a Markov Chain Monte Carlo (MCMC) algorithm that can explore all possible configurations that are compatible with the data. The model and the associated MCMC algorithm are validated on synthetic data and then applied to a subsoil in the Venetian Plain with a moderately dense network of cored boreholes.

**Keywords** Subsoil modeling · Stratigraphic sequence · PC prior · Stochastic 3D model · Data augmentation · Conditional simulation

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## 1 Introduction

The case study motivating this work is a subsoil in the Venetian Plain with a moderately dense network of cored boreholes. Geologists and hydrogeologists managing this subsoil are in need of stochastic three-dimensional models of the stratigraphic sequence. The model should of course be conditioned to borehole data. The sequence of layers must correspond to the known regional stratigraphic sequence and, in addition, to the surfaces separating the layers are required to be smooth and continuous.

Simulating a depositional (or stratigraphic) sequence conditionally on boreholes data has been and still is a long-standing problem in hydrogeology and in petroleum geostatistics. In the context of reservoir modeling, Pyrcz et al. (2015) offers a comprehensive overview of the literature and a convincing conceptual framework in which methods are represented along a complexity gradient with one extreme corresponding to pixel based models with statistics and conditioning derived from the data and the other extreme representing geological concepts unconditional to local observations. As models tend to move away from the less complex extreme to the more complex one, they are less versatile and more difficult to condition (Pyrcz et al., 2015). Easy-to-condition pixel based methods thus tend to be favored when data are dense, whereas rule-based or process-based models are preferred when conditioning data is sparse.

Pixel based approaches, whether based on variograms (Matheron et al., 1987), truncated Gaussian random fields and plurigaussian random fields (Beucher et al., 1993; Galli et al., 1994; Armstrong et al., 2011; Le Blévec et al., 2017, 2018), transiograms (Carle and Fogg, 1996), or MCP (Allard et al., 2011; Sartore et al., 2016; Benoit et al., 2018b), are well known and relatively easy to handle. For these approaches, variogram and transiogram fitting is well understood and conditioning to well data is efficient, even for truncated Gaussian models (Marcotte and Allard, 2018). However, one source of difficulty in the fitting procedure is the fact that the processes and the amount of information are often anisotropic. Typically, for borehole data, there is much more information along the depth than along horizontal directions.

Multiple point statistics (MPS) approaches (Strebel, 2002; Mariethoz and Caers, 2014) require a training image when simulations are performed in two dimensions. Three-dimensional simulations are much more difficult to perform, since training cubes are rarely available at kilometer scales. Methods for combining images in three-dimensional simulations have been proposed (Comunian et al., 2012, 2014). But since a high degree of continuity is required for layers in this work, pixel based methods, including MPS, are not deemed appropriate.

Object models, such as Boolean models, are more difficult to fit and to condition, in particular when accounting for non-stationarity and erosion rules, see for example Syversveen and Omre (1997) and Allard et al. (2006). In addition, object models are not geologically appropriate for simulating sequences of layers.

46 Rule-based and process-based models incorporate some amount of under-  
47 standing of the geological processes. They use rules to control the temporal  
48 sequence and spatial position of geological objects so as to mimic geological  
49 processes. Among others, they have been applied to fluvial systems, deepwater  
50 channel systems and turbiditic lobes systems. Particular cases of interest to  
51 this work are surface-based models. For simulating lobes in a turbidite reser-  
52 voir, Bertoncello et al. (2013) proposed a rule-based stacking of lobe-shape  
53 events with quite complicated sequential placement rules that depend partly  
54 on the already simulated events. The conditioning to well-log data and seis-  
55 mic data is achieved through sequential optimization. One of the limitations  
56 of this approach is that the variability between the conditional simulations is  
57 low, owing to the optimization approach. A second limitation recognized by  
58 the authors is that their method works best with a limited amount of data.

59 This paper presents a new rule-based approach for simulating depositional  
60 sequences of surfaces conditionally to lithofacies thickness data. It is a stochas-  
61 tic model that belongs to the *Markov rules* sub-class of rule-based methods, see  
62 Pyrcz et al. (2015) and appropriate references therein. The thickness of each  
63 layer is modeled by a transformed latent Gaussian random field allowing for  
64 null thickness. The random fields are *latent* because they can be unobserved on  
65 some parts of the domain under study, thanks to a truncation process. Layers  
66 are sequentially stacked above each other following the regional stratigraphic  
67 sequence. By choosing adequately the variograms of these random fields, the  
68 simulated surfaces separating two layers can be continuous and smooth. Con-  
69 ditioning to the observed borehole data is made possible thanks to constrained  
70 Gaussian conditioning, as will be explained later on.

71 A problem that has been barely addressed in geostatistical models for de-  
72 positional sequences is the fact that borehole information is often incomplete  
73 in the sense that it does not provide direct information regarding the exact  
74 layers that have been observed. For example, let us consider that the strati-  
75 graphic sequence of the study domain contains several repetitions of a given  
76 lithofacies, say Clay. Consider also that the recorded data at one given bore-  
77 hole measures one single thickness for Clay. A first possibility is that there is  
78 actually only one Clay layer at this location, but it could be any of the several  
79 Clay layers of the regional stratigraphic sequence. Simulations should there-  
80 fore account for this uncertainty. A second possibility is that the measurement  
81 actually corresponds to two (or more) Clay layers, one on top of the other,  
82 with missing intermediate layers at this location. In this case, the measured  
83 thickness should be shared between two layers. The latent Gaussian model  
84 proposed in this paper offers a natural solution to this problem by means of  
85 a Bayesian setting with a Markov Chain Monte Carlo (MCMC) algorithm  
86 that can explore all possible configurations compatible with the data. Notice  
87 that the approach proposed in Bertoncello et al. (2013) does not address this  
88 problem at all.

89 The rest of this paper is organized as follows. Section 2 is devoted to the  
90 conceptual model. In particular, the difference between the (unique) regional  
91 stratigraphic sequence, referred to as the parent sequence, and the observed

92 sequences is detailed. Section 3 presents the stochastic model. In Sect. 4 all  
 93 details for Bayesian inference with an MCMC algorithm are given. It is then  
 94 validated on a synthetic data set in Sect. 5. Finally, it is successfully applied  
 95 to the Venetian Plain that motivated this work in Sect. 6. Some concluding  
 96 remarks are then given in Sect. 7.

## 97 2 The Conceptual Model

### 98 2.1 Notations

99 Let us consider a spatial domain  $\mathcal{S} \in \mathbb{R}^2$  and an interval  $\mathcal{T} \subset \mathbb{R}^+$ , which will  
 100 correspond to “depth”. Note that depth can be converted into time through  
 101 depositional processes, which is the reason why  $t \in \mathcal{T}$  is used to denote depth.  
 102 Let us also consider a family of  $K$  lithofacies,  $\mathcal{C} = \{C_1, \dots, C_K\}$ . The aim of  
 103 this work is to build a process  $X = \{X(s, t)\}$ , defined at any point  $(s, t) \in \mathcal{S} \times \mathcal{T}$   
 104 and taking values in  $\mathcal{C}$ . In other words, at each location is associated one and  
 105 only one lithofacies. The process must be continuous almost everywhere and  
 106 the discontinuity surfaces should be smooth and have a general horizontal  
 107 orientation. The process  $X$  is observed along depth at a finite number of  
 108 locations  $s_1, \dots, s_n$  and each observation corresponds to a drilled core, referred  
 109 to as boreholes in the rest of this work.

110 Let  $X_i = \{X(s_i, t), t \in \mathcal{T}\}$  be one of these observations at site  $s_i$ ,  $i =$   
 111  $1, \dots, n$ , where  $n$  is the number of sites. The observation  $X_i$  is piece-wise  
 112 constant, with  $M_i$  discontinuities at different depths each time a new layer  
 113 is encountered. The resulting information is a sequence of facies and depths,  
 114 referred to as the *observed sequence*,  $(\mathbf{C}_i^o, \mathbf{T}_i^o)$ , where  $\mathbf{C}_i^o = (C_{1,i}^o, \dots, C_{M_i,i}^o)$   
 115 with  $C_{j,i}^o \in \mathcal{C}$  for  $j = 1, \dots, M_i$ , and  $\mathbf{T}_i^o = (T_{1,i}^o, \dots, T_{M_i,i}^o)$  with  $T_{j,i}^o \in \mathcal{T}$   
 116 and  $T_{1,i}^o < \dots < T_{M_i,i}^o$ . The depths are measured with respect to a ground-  
 117 level  $T_{0,i}$ . The thicknesses of each observed layer  $\mathbf{Z}_i^o = (Z_{1,i}^o, \dots, Z_{M_i,i}^o)$  can  
 118 be derived from the depths, with  $Z_{j,i}^o = T_{j,i}^o - T_{j-1,i}^o$ ,  $j = 1, \dots, M_i$ . Finally,  
 119 the last layer is assumed to be completely observed, that is the depth  $Z_{M_i,i}^o$  is  
 120 assumed to be not censored.

### 121 2.2 Parent Sequence

122 The working hypothesis is that there exists a common lithological sequence  
 123 of facies, hereafter referred to as the “parent sequence,” which is compatible  
 124 with all observed sequences in the area of study in the sense that each observed  
 125 sequence can be obtained from the parent sequence by deleting some layers of  
 126 the parent sequence.

127 This sequence can result from the prior knowledge of the scientists. Al-  
 128 ternatively, it can be derived from the observed data. From a mathematical  
 129 viewpoint, there always exists a parent sequence. For example, it can easily  
 130 be obtained by simply stacking all observed sequences into a single sequence.

131 Then, each observed sequence of layers is simply obtained by “deleting” all  
 132 other observed sequences. Obviously, this parent sequence is of no modeling  
 133 interest, but it is mathematically important since it provides a proof of the ex-  
 134 istence of this concept. In general, very long parent sequences are uninteresting  
 135 from a modeling point of view. In accordance with a parsimony principle, one  
 136 should seek the shortest possible parent sequences. Clearly, there is only a finite  
 137 number of parent sequences of minimal length. Such parent sequences could  
 138 be built using discrete optimization algorithms, or they could be provided by  
 139 scientists, based on prior geological knowledge. Either way, how minimal par-  
 140 ent sequences are obtained is a subject out of the scope of the present research,  
 141 and this route is not pursued any longer.

142 From now on it will be considered that the parent sequence is known,  
 143 and that it is one of the minimal length parent sequences. The parent se-  
 144 quence of length  $M$  will be denoted  $\mathbf{C} = (C_1, \dots, C_M)$ ,  $C_i \in \mathcal{C}$ , with  $M \geq$   
 145  $\max\{K, M_1, \dots, M_n\}$ .

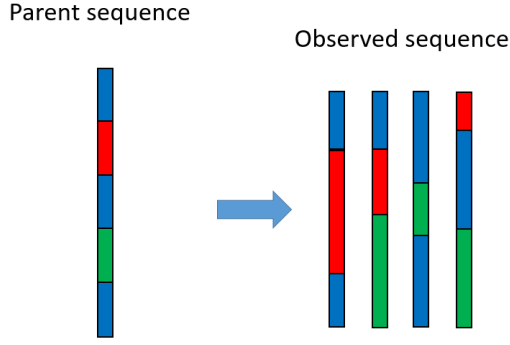
### 146 2.3 From the Parent Sequence to the Observed Sequences

147 When analyzing sequences of lithofacies, it is quite common that some facies  
 148 are unobserved at one or several boreholes. In order to allow for this, each  
 149 observed sequence at each site  $s_i$  is therefore a subset of a complete sequence  
 150  $(\mathbf{C}, \mathbf{T}_i)$  corresponding to the parent sequence. The corresponding vector of  
 151 complete thickness is  $\mathbf{Z}_i$ , and, in contrast to the observed ones, some thickness  
 152  $Z_{j,i} = T_{j,i} - T_{j-1,i}$ ,  $j = 1, \dots, M$  can be equal to zero. In this case, the  
 153 corresponding layer is unobserved at location  $s_i$ . When  $M_i < M$ , the sequence  
 154 at  $s_i$  is an *incomplete sequence*, and  $\mathbf{C}_i^o$  is a sub-sequence of  $\mathbf{C}$ . The complete  
 155 data will be denoted  $\mathbf{X} = \{(\mathbf{C}, \mathbf{Z}_i), i = 1, \dots, n\}$  and  $\mathbf{X}^o = \{(\mathbf{C}_i^o, \mathbf{Z}_i^o), i =$   
 156  $1, \dots, n\}$  will denote the observed data. In the following,  $O(\cdot)$  will denote the  
 157 mapping such that  $\mathbf{X}^o = O(\mathbf{X})$ . Figure 1 illustrates a parent sequence and  
 158 four different possible observed sequences.

## 159 3 Statistical Setting

### 160 3.1 Stochastic Model

161 The stochastic model requires a univariate model for the marginal distribution  
 162 of the thicknesses and a spatial model to account for the lateral continuity of  
 163 the layers. Thicknesses are modeled using positive zero inflated random vari-  
 164 ables in order to account for the many 0s resulting from incomplete observed se-  
 165 quences. Among many possible models, latent truncated Gaussian models (All-  
 166 croft and Glasbey, 2003; Baxevani and Lennartsson, 2015; Benoit et al., 2018a),  
 167 also referred to as Tobit models (Liu et al., 2019) in econometrics, are flexible  
 168 models that easily allow geostatistical modeling. Spatial dependence among  
 169 the thicknesses belonging to a same layer is introduced by means of a trun-  
 170 cated Gaussian random field. More precisely, for  $j = 1, \dots, M$ , let  $W_j(s)$ ,  $s \in \mathcal{S}$



**Fig. 1** Parent sequence and four possible incomplete observed sequences. Since the parent sequence is conceptual, thicknesses are only meaningful in the observed sequences.

171 be a standardized Gaussian random field that, for simplicity, will be supposed  
 172 stationary with covariance function  $\text{cov}[W_j(s), W_j(s')] = \rho_j(s - s'; \xi_j)$ , where  
 173  $\rho_j$  is a parametric correlation function and  $\xi_j$  the vector of associated param-  
 174 eters. The thickness field  $\{Z_j(s), s \in \mathcal{S}\}$  is defined as

$$Z_j(s) = \varphi_j(W_j(s) - \tau_j) \quad \text{if } W_j(s) > \tau_j, \quad (1)$$

175 and  $Z_j(s) = 0$  otherwise, where  $\tau_j$  is a threshold and  $\varphi_j(\cdot)$  is a continuous  
 176 one-to-one mapping from  $\mathbb{R}_+$  to  $\mathbb{R}_+$ . The probability of positive thickness  
 177  $\Pr(Z_j(s) > 0)$  will be denoted by  $p_j$ . With this construction, null thickness  
 178 has a positive probability, since  $\Pr(Z_j(s) = 0) = 1 - p_j = \Phi(\tau_j) > 0$ , where  
 179  $\Phi(\cdot)$  is the cumulative probability function of the standard Gaussian random  
 180 variable. Parameters of the stochastic model can be expressed equivalently in  
 181 terms of  $\tau_j$  or  $p_j$  and in the sequel the second setting is chosen. One particular  
 182 case that will be used later is to set  $\varphi_j(x) = \mu_j x^{\beta_j}$ ,  $x > 0$  with  $\beta_j, \mu_j > 0$ .  
 183 When  $\beta_j = 1$ , one gets

$$E[Z_j(s)] = \mu_j \left( \frac{\phi(\tau_j)}{1 - \Phi(\tau_j)} - \tau_j \right), \quad (2)$$

184 and

$$\text{Var}[Z_j(s)] = \mu_j^2 \left[ 1 + \frac{\phi(\tau_j)}{1 - \Phi(\tau_j)} \left( \tau_j - \frac{\phi(\tau_j)}{1 - \Phi(\tau_j)} \right) \right], \quad (3)$$

185 where  $\phi(\cdot)$  is the density function of the standard Gaussian random variable.  
 186 When  $\beta_j$  is not an integer, the moments of  $Z_j(s)$  involve hypergeometric func-  
 187 tions and are not reported here. From Eqs. (2) and (3), it is clear that the  
 188 expectation and standard deviation of the thickness of layer  $j$  are both propor-  
 189 tional to the parameter  $\mu_j$ . The covariance function  $\rho_j$  must be smooth enough  
 190 in order to generate regular thicknesses. For example, choosing that  $\rho_j$  is twice  
 191 differentiable at the origin leads to a mean-squared differentiable random field  
 192  $W_j$  and, as a consequence, to a mean-squared differentiable random field for  
 193 the thicknesses since  $\varphi_j$  is continuous and locally finite boundaries of the non

194 null thickness sets. The depth surfaces  $\{T_j(s), s \in \mathcal{S}\}$  are then obtained by  
 195 adding up the thickness fields. Starting from a fixed and known ground-floor  
 196  $T_0 = \{T_0(s), s \in \mathcal{S}\}$  one sets

$$T_j(s) = T_{j-1}(s) + Z_j(s) = T_0(s) + \sum_{i=1}^j Z_i(s), \quad j = 1, \dots, M.$$

197 Finally, the random fields  $W_j$  are assumed to be independent, since they relate  
 198 to independent depositional processes.

### 199 3.2 Complete Likelihood

200 Since layers are assumed to be independent, the complete likelihood factorizes  
 201 into a product of  $M$  likelihoods

$$L(\theta; \mathbf{X}) = \prod_{j=1}^M L_j(\theta_j; Z_{j,1}, \dots, Z_{j,n}), \quad (4)$$

202 where  $\theta_j = (p_j, \mu_j, \beta_j, \xi_j), j = 1, \dots, M$  and  $\theta = (\theta_1, \dots, \theta_M)$ .

203 In the sequel  $\phi_k(\cdot, \boldsymbol{\mu}, \boldsymbol{\Sigma})$  and  $\Phi_k(\cdot, \boldsymbol{\mu}, \boldsymbol{\Sigma})$  denote the density and the cu-  
 204 mulative distribution function of a  $k$ -multivariate Gaussian random variable  
 205 with mean vector  $\boldsymbol{\mu}$  and covariance matrix  $\boldsymbol{\Sigma}$ . Let us consider now a layer  
 206  $j \in \{1, \dots, M\}$ . For convenience, thicknesses and the corresponding locations  
 207 are reordered such that the first  $n_j$  thicknesses  $Z_{j,1}, \dots, Z_{j,n_j}$  correspond to  
 208 the positive values and the remaining  $\ell_j = n - n_j$  ones are 0. The complete-data  
 209 likelihood of the single layer  $j$  is

$$L_j(\theta_j; Z_{j,1}, \dots, Z_{j,n}) = f_j(Z_{j,1}, \dots, Z_{j,n_j}; \theta_j) F_j(0, \dots, 0, |Z_{j,1}, \dots, Z_{j,n_j}; \theta_j). \quad (5)$$

210 The density  $f_j(Z_{j,1}, \dots, Z_{j,n_j}; \theta_j)$  is given by

$$f(Z_{j,1}, \dots, Z_{j,n_j}; \theta) = \phi_{n_j}(W_{j,1}, \dots, W_{j,n_j}; \mathbf{0}, \boldsymbol{\Sigma}_j) \prod_{i=1}^{n_j} J_{\varphi_j^{-1}}(Z_{j,i}), \quad (6)$$

211 where  $\boldsymbol{\Sigma}_j = \boldsymbol{\Sigma}_{n_j, n_j} = [\rho(s_i - s_k; \xi_j)]_{i,k=1, \dots, n_j}$ ,  $W_{j,i} = \varphi_j^{-1}(Z_{j,i}) + \tau_j$ ,  $i =$   
 212  $1, \dots, n_j$ , and  $J_{\varphi_j^{-1}}(Z_{j,i})$  is the Jacobian of  $\varphi_j^{-1}$  computed at  $Z_{j,i}$ . The condi-  
 213 tional probability  $F_j(0, \dots, 0 | Z_{j,1}, \dots, Z_{j,n_j}; \theta)$  is given by

$$F_j(0, \dots, 0 | Z_{j,1}, \dots, Z_{j,n_j}; \theta) = \Phi_{\ell_j}(\tau_j, \dots, \tau_j; \mathbf{m}_j, \mathbf{V}_j), \quad (7)$$

214 where the mean vector  $\mathbf{m}_j$  and covariance matrix  $\mathbf{V}_j$  can be easily derived  
 215 using the kriging equations (Cressie, 1993; Chilès and Delfiner, 2012)

$$\mathbf{m}_j = \boldsymbol{\Sigma}_{\ell_j, n_j} \boldsymbol{\Sigma}_{n_j, n_j}^{-1} \mathbf{W}_{n_j}; \quad \mathbf{V}_j = \boldsymbol{\Sigma}_{\ell_j, \ell_j} - \boldsymbol{\Sigma}_{\ell_j, n_j} \boldsymbol{\Sigma}_{n_j, n_j}^{-1} \boldsymbol{\Sigma}_{n_j, \ell_j}, \quad (8)$$

with  $\mathbf{W}_{n_j} = (W_{j,1}, \dots, W_{j,n_j})'$  and the matrices  $\Sigma_{\ell_j, n_j}$  and  $\Sigma_{\ell_j, \ell_j}$  being defined in similar ways as  $\Sigma_{n_j, n_j}$ . To summarize, the complete data likelihood in (4) becomes

$$\begin{aligned} L(\theta; \mathbf{X}) &= \prod_{j=1}^M L_j(\theta; Z_{j,1}, \dots, Z_{j,n}) \\ &= \prod_{j=1}^M \phi_{n_j}(W_{j,1}, \dots, W_{j,n}; \mathbf{0}, \Sigma_j) \\ &\quad \times \prod_{i=1}^{n_j} J_{\varphi_j^{-1}}(Z_{j,i}) \Phi_{l_j}(\tau_j, \dots, \tau_j; \mathbf{m}_j, \mathbf{V}_j). \end{aligned} \quad (9)$$

In the particular case  $\varphi_j(x) = \mu_j x^{\beta_j}$  that will be considered below, the Jacobian simplifies to

$$J_{\varphi_j^{-1}}(Z_{j,i}) = \frac{1}{\mu_j \beta_j} \left( \frac{Z_{j,i}}{\mu_j} \right)^{1-1/\beta_j}. \quad (10)$$

### 3.3 Observed Likelihood

In principle, the observed likelihood is related to the complete likelihood through

$$L(\theta; \mathbf{X}^o) = \int_{\{\mathbf{X}: \mathbf{X}^o = O(\mathbf{X})\}} L_X(\theta; \mathbf{X}) d\mathbf{X}. \quad (11)$$

However, even for moderately long parent sequence and number of 0 thicknesses, the space  $\{\mathbf{X} : \mathbf{X}^o = O(\mathbf{X})\}$  is difficult to explore and the integral (11) becomes intractable. These difficulties are illustrated with two examples. At some site, let us consider an observed sequence  $(\mathbf{C}^o, \mathbf{T}^o)$  and the corresponding thicknesses  $\mathbf{Z}^o$ . Here, the reference to the site is dropped for the sake of clearer notations. Recall that since the sequence  $\mathbf{C}^o$  must be compatible with the parent sequence  $\mathbf{C}$ ,  $\mathbf{C}^o$  is obtained by deleting some layers of  $\mathbf{C}$ .

Table 1 shows an example of a parent sequence  $\mathbf{C}$  with three categories: **Blue**, **Red** and **Green**. The observed sequence  $\mathbf{C}^o$  is incomplete. Several augmented sequences  $\mathbf{C}^a$  with corresponding depths  $\mathbf{T}^a$  are possible. Since in the observed series the first **Blue** is followed by **Red**, the sub-sequence **[Blue-Red]** must correspond to the beginning of the parent sequence. Regarding the second occurrence of **Blue**, three cases can be distinguished: i) it corresponds only to the third layer of  $\mathbf{C}$  with 4th and 5th layers having null thickness; ii) it corresponds only to the fifth layer, in which case the 3rd and 4th layers have null thickness; iii) it corresponds partly to the 3rd and partly to the 5th layers. Then, only the 4th layer has 0 thickness. In this last case, an intermediate, latent, transition at depth  $\tilde{T}$  with  $T_2^o \leq \tilde{T} \leq T_3^o$  must be introduced. These augmented series are all possible, but some will be more likely than others, depending on the parameters of the model. In Appendix A an even more complex example is provided. Only some of the possible configurations are shown.



245 They are too numerous and complex to be completely listed, even for short  
246 parent sequences.

247 In order to estimate the parameters of the model, a data augmentation  
248 algorithm (Tanner, 1996, Ch. 5) can be exploited where the complete sequences  
249 that are compatible with the observed ones are explored. A Bayesian approach  
250 will be adopted for the inference of the parameters and a Markov Chain Monte  
251 Carlo (MCMC) algorithm will be designed in Sect. 4. But first, a simulation  
252 in which all parameters are known and all sequences are complete is shown.

**Table 1** Example of a parent sequence  $\mathbf{C}$  with an observed sequence  $\mathbf{C}^o$  and several possible augmented sequences with corresponding transition depths and thicknesses

Parent	Observed		Possible augmented sequences								
$\mathbf{C}$	$\mathbf{C}^o$	$\mathbf{T}^o$	$\mathbf{C}^a$	$\mathbf{T}^a$	$\mathbf{Z}^a$	$\mathbf{C}^a$	$\mathbf{T}^a$	$\mathbf{Z}^a$	$\mathbf{C}^a$	$\mathbf{T}^a$	$\mathbf{Z}^a$
Blue	Blue	$T_1^o$	Blue	$T_1^o$	$T_1^o$	Blue	$T_1^o$	$T_1^o$	Blue	$T_1^o$	$T_1^o$
Red	Red	$T_2^o$	Red	$T_2^o$	$T_2^o - T_1^o$	Red	$T_2^o$	$T_2^o - T_1^o$	Red	$T_2^o$	$T_2^o - T_1^o$
Blue	Blue	$T_3^o$	Blue	$T_3^o$	$T_3^o - T_2^o$		$T_2^o$	0	Red	$\tilde{T}$	$\tilde{T} - T_2^o$
Green	-	-		$T_3^o$	0		$T_2^o$	0		$\tilde{T}$	0
Blue	-	-		$T_3^o$	0	Blue	$T_3^o$	$T_3^o - T_2^o$	Blue	$T_3^o$	$T_3^o - \tilde{T}$

### 253 3.4 Simulation

254 Unconditional simulation is straightforward when the transformation  $\varphi_j$  and  
255 the parameters  $\theta_j$ ,  $j = 1, \dots, M$ , are known. All that is required is to sim-  
256 ulate  $M$  random fields  $W_j$ ,  $j = 1, \dots, M$  and then to apply (1) in order to  
257 transform the Gaussian process into a thickness surface. Figure 2 illustrates a  
258 cross-section of a two-dimensional simulation over  $\mathcal{S} = [0, 100] \times [0, 100]$  with  
259 four lithofacies {**Black-Red-Blue-Green**} and  $\varphi_j(x) = \mu_j x$ , that is  $\beta_j = 1$  for  
260 all categories. The parent sequence has 15 layers (see Fig. 2-(a)) and stochastic  
261 models for layers with the same lithofacies have identical parameters. Thick-  
262 nesses have been simulated using Gaussian random fields with a Matérn co-  
263 variance function

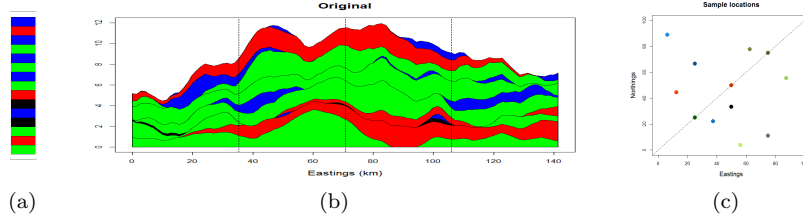
$$\rho(h; \nu, \alpha, \sigma^2) = \frac{\sigma^2}{2^{\nu-1} \Gamma(\nu)} \left( \frac{\|h\|}{\alpha} \right)^\nu K_\nu \left( \frac{\|h\|}{\alpha} \right), \quad h \in \mathbb{R}^2, \quad (12)$$

264 where  $\nu > 0$  is a smoothness parameter,  $\alpha > 0$  a range parameter and  $\sigma^2$   
265 the sill.  $\Gamma$  is the gamma function and  $K_\nu$  is the modified Bessel function of  
266 the second kind of order  $\nu$ . Here, the smoothness parameter has been set to  
267  $\nu = 3/2$  and  $\sigma^2 = 1$ , which leads to the simplified expression  $\rho_j(h; \alpha_j) =$   
268  $(1 + \|h\|/\alpha_j) \exp(-\|h\|/\alpha_j)$ , where  $\alpha_j$  is a range parameter. The set of the  
269 parameters in the simulation experiment is shown in Table 2.

270 Twelve synthetic boreholes have been located in  $\mathcal{S}$ . Three of them are  
271 placed along the diagonal at coordinates (25, 25), (50, 50) and (75, 75). Nine

**Table 2** Parameters for the simulation example.

	Black	Red	Blue	Green
$\mu, \beta$	1	1	1	1
$p$	0.3	0.8	0.3	0.8
$\alpha$	20	20	10	10

**Fig. 2** Simulation experiment: (a) Parent sequence of length 15, with 4 lithofacies {Black-Red-Blue-Green}; (b) Cross-section of a two-dimensional simulation along the diagonal of  $S = [0, 100] \times [0, 100]$ . See Table 2 for the parameters; (c) Locations of the twelve boreholes.

272 others are randomly located (see Fig. 2(c)). For each category, the observed  
 273 frequencies along these twelve boreholes are (0.58, 0.83, 0.28, 0.80). Notice that  
 274 **Black** is highly over-represented. The average thicknesses computed along the  
 275 boreholes are (0.31, 1.31, 0.62, 1.25) for each of the four categories, whilst the  
 276 theoretical expectations of each category computed as per (2) are respectively  
 277 (1.8, 1.1, 1.8, 1.1). Note here that **Black** and **Blue** are very unlikely to be di-  
 278 rectly stacked above each other, while it is often the case for **Red** and **Green**.

279 Conditional simulation is relatively easy to implement when the parameters  
 280 are known and when complete sequences of thicknesses are available, including  
 281 all null thicknesses. Care must be taken when simulating values from the Gaus-  
 282 sian distribution that are below the thresholds  $\tau_j$ , but otherwise the algorithm,  
 283 shown in Algorithm 1, is rather straightforward. Simulations of the truncated  
 284 Gaussian values are done by calling the function `rmvnorm` of the R package  
 285 `mvtnorm` (Genz et al., 2019). The reader is referred to Chilès and Delfiner  
 286 (2012) for a general exposition on unconditional simulations and conditional  
 287 simulations using Kriging techniques.

## 288 4 Bayesian Inference with a Markov Chain Monte Carlo Algorithm

### 289 4.1 Sampling All Possible Configurations

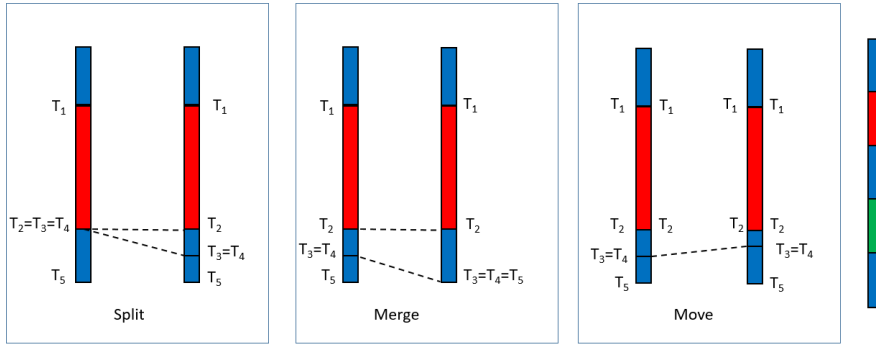
290 In order to sample within all possible configurations of the augmented sequence  
 291 at a given site  $s_i$  that are compatible with the parent sequence, the Markov  
 292 Chain Monte Carlo (MCMC) algorithm must be able to delete a layer, to add  
 293 a new layer or to displace the limit between two layers of the same category.  
 294 Recall that the limit between two different categories are hard conditioning

**Algorithm 1** Conditional simulation when all sequences and all parameters are known

**Require:** Data with complete sequences; transform functions  $\varphi_j, j = 1, \dots, M$

**Require:** All parameters

- 1: **for**  $j = 1$  to  $M$  **do**
- 2:   Compute the vector  $\mathbf{W}_{n_j} = (W_{j,1}, \dots, W_{j,n_j})$  where  $W_{j,k} = \varphi_j^{-1}(Z_{j,k})$  corresponding to  $Z_{j,k} > 0, k = 1, \dots, n_j$
- 3:   Compute  $\mathbf{m}_j$  and  $\mathbf{V}_j$  according to (8)
- 4:   Draw a vector of length  $\ell_j$  from a truncated multivariate Gaussian distribution,  $\mathbf{W}_{\ell_j} \sim \mathcal{TN}_{\ell_j}(\mathbf{m}_j, \mathbf{V}_j; -\infty, \tau_j)$ , for which each component must be below  $\tau_j$ .
- 5:   Set  $\mathbf{W}_j = (\mathbf{W}_{n_j}, \mathbf{W}_{\ell_j})$
- 6:   Simulate a Gaussian random field  $F_j$  conditionally on  $\mathbf{W}_j$
- 7:   Transform the field  $F_j$  into the thicknesses according to (1)
- 8: **end for**



**Fig. 3** Elementary moves in an incomplete observed sequence. Note that the layer **Green** is unobserved. From left to right: *Split*, *Merge* and *Displace*.

295 data that cannot be changed. These elementary moves, illustrated in Fig. 3,  
 296 are now detailed.

297 *Split*: A state is split into two successive states of the same category. A split  
 298 is only possible if it is compatible within the parent sequence. For example,  
 299 in Fig. 3, the **Blue** layer at the bottom can be split into two layers since  
 300 the parent sequence contains a second **Blue** layer. In Table 5 the situation  
 301 in panel 4 can be obtained by splitting the state **Red**, either in panel 2  
 302 or in panel 3. When a state is split, a new transition depth, denoted  $t_i$   
 303 in Table 5, must be introduced. The thickness is split in two thicknesses  
 304 accordingly.

305 *Merge*: This move is the opposite move of *Split*. Two successive states in the  
 306 same category are merged together. The corresponding depth is removed  
 307 and the resulting thickness is the sum of the two merged thicknesses.

308 *Displace*: Here, the augmented sequence is not changed, but the intermediate  
 309 value between two successive states of the same category is changed. The  
 310 corresponding thicknesses are then updated.

311 It is easy to verify that, starting from any initial configuration that is  
 312 compatible with the parent sequence, any other configuration can be reached  
 313 by combining finite numbers of *Split*, *Merge* and *Displace*. Hence, if these  
 314 moves are used as building blocks of a MCMC algorithm, the resulting Markov  
 315 Chain will be ergodic. At each borehole, one of the three moves is proposed  
 316 with probabilities  $(p_S, p_M, p_D)$  with  $p_S + p_M + p_D = 1$ . If the move is possible,  
 317 it is accepted according to Metropolis-Hasting acceptance ratio described in  
 318 Sect. 4.3.

## 319 4.2 Choosing the Priors

320 Priors must be defined for all parameters of the model. For the parameters of  
 321 the transform functions  $\varphi_j$ ,  $1 - p_j = \Phi(\tau_j)$  and  $\beta_j$ , uninformative flat priors  
 322 have been chosen on the intervals  $(0, 1)$  and  $(0.25, 4)$ , respectively. Regarding  
 323 the covariance function, the Matérn covariance function in (12) has been  
 324 chosen for its great flexibility thanks to three parameters:  $\xi = (\nu, \alpha, \sigma)$ , for  
 325 smoothness, range and sill, respectively.

326 However, it is known that the joint estimation of these parameters is diffi-  
 327 cult in a Bayesian context, in particular if the number of data is small. Zhang  
 328 (2004) showed that for a Matérn covariance function the only quantity that  
 329 can be estimated consistently under in-fill asymptotics is  $\sigma^2 \alpha^{-2\nu}$ . As a conse-  
 330 quence, since the parameter  $\mu^2$  behaves as the marginal variance of the random  
 331 field, using uninformative flat priors for  $(\mu, \alpha, \nu)$  is expected to provide poor  
 332 posterior distributions for these parameters. This was indeed confirmed on  
 333 preliminary MCMC runs (results not reported here). It was thus decided to  
 334 fix the smoothness parameter  $\nu$  among the values  $(1/2, 3/2, 5/2)$  that would  
 335 provide the highest likelihood. The above values correspond to covariance func-  
 336 tions being the product of an exponential and a polynomial of order  $p$  with  
 337  $p = 0, 1, 2$  respectively, namely  $\rho(r; 1/2, \alpha, 1) = \exp(-r/\alpha)$ ,  $\rho(r; 3/2, \alpha, 1) =$   
 338  $(1 + r/\alpha) \exp(-r/\alpha)$  and  $\rho(r; 5/2, \alpha, 1) = (1 + r/\alpha + r^2/(3\alpha^2)) \exp(-r/\alpha)$ .

339 Simpson et al. (2017) proposed an approach for building priors that are  
 340 based on penalizing the complexity to a base model. For example, a random  
 341 effect with positive variance is an extension (a more complex version) of ran-  
 342 dom effect with null variance. Similarly, a random field with a finite range is  
 343 an extension (a more complex version) of a random field with an infinite range.  
 344 Indeed, if the range is infinite, the random field is perfectly correlated and its  
 345 spatial variance is null. Penalized Complexity (PC) priors are then defined as  
 346 the only priors that: i) use the Kullback-Leibler divergence as a measure be-  
 347 tween the extended and the base models; ii) have a penalization that increases  
 348 with the distance at a constant rate.

349 Fuglstad et al. (2019) derived the PC priors for a Matérn covariance with  
 350 parameters  $\sigma$ ,  $\alpha$  and  $\nu$ , when  $\nu$  is fixed. They showed that the joint PC prior  
 351 corresponding to a base model with infinite range and zero variance when  
 352  $d = 2$  is

$$\pi(\sigma, \alpha) = \lambda_\alpha \alpha^{-2} \exp(-\lambda_\alpha/\alpha) \lambda_\sigma \exp(-\lambda_\sigma \sigma), \quad (13)$$

353 where  $\lambda_\alpha = -\ln(\epsilon_\alpha)\alpha_0$  and  $\lambda_\sigma = -\ln(\epsilon_\sigma)/\sigma_0$ , and the values of  $\lambda_\alpha$  and  $\lambda_\sigma$  are  
 354 such that  $P(\alpha < \alpha_0) = \epsilon_\alpha$  and  $P(\sigma > \sigma_0) = \epsilon_\sigma$ . By choosing small probabilities  
 355  $\epsilon_\alpha$  and  $\epsilon_\sigma$ , the range is lower-bounded above  $\alpha_0$  and the standard deviation  
 356 is upper bounded at  $\sigma_0$  with probability  $1 - \epsilon_\alpha$  and  $1 - \epsilon_\sigma$ , respectively. PC  
 357 priors described in (13) will be used throughout, where  $\mu$  plays the role of the  
 358 standard deviation as shown in Eq. (3) in Sect. 3.1.

### 359 4.3 General Description of the Algorithm

360 Each parameter in each category is updated iteratively in a Metropolis-within-  
 361 Gibbs algorithm (Gelfand, 2000). A new value is proposed according to sym-  
 362 metric transition kernels, for which it is equally likely to move from a current  
 363 value  $y^c$  to a new value  $y^n$  than the opposite. Let  $\theta^c$  and  $\theta^n$  be the current  
 364 and the proposed vector of parameters  $\theta$ , respectively. Let further  $\pi(\cdot)$  be the  
 365 prior density of  $\theta$ . The acceptance ratio is then

$$A(\theta^c, \theta^n) = \frac{L(\theta^n; \mathbf{X})\pi(\theta^n)}{L(\theta^c; \mathbf{X})\pi(\theta^c)}. \quad (14)$$

366 When sampling the configurations thanks to one of the possible moves *Split*,  
 367 *Merge* and *Displace*, a new configuration  $\mathbf{X}^n$  is proposed,  $\mathbf{X}^c$  being the current  
 368 one. In this case, the acceptance ratio is

$$A(\mathbf{X}^c, \mathbf{X}^n) = \frac{L(\theta; \mathbf{X}^n)}{L(\theta; \mathbf{X}^c)}. \quad (15)$$

369 The proposals are accepted if the acceptance ratios  $A(\cdot, \cdot)$  are larger than  
 370 one. Otherwise, they are accepted with a probability equal to the ratio. The  
 371 proposal in the Metropolis-Hasting step are random walk proposals aiming  
 372 at an acceptance rate above 0.5. For sampling new configurations at each  
 373 borehole in turn, a possible move is drawn according to the probabilities  $p_S =$   
 374  $p_M = p_D = 1/3$ . Then, it is checked whether such a move is feasible within  
 375 this borehole. If several moves are possible, one is selected uniformly among  
 376 all possible moves in that borehole, and a new configuration is proposed. The  
 377 whole procedure is summarized in Algorithm 2.

## 378 5 A Synthetic Data Example

379 The MCMC algorithm described above is first validated on the synthetic data-  
 380 set described in Sect. 3.4 and illustrated in Fig. 2. It was coded in R using stan-  
 381 dard functions and our own code for the *Split*, *Merge* and *Displace* movements.  
 382 Most of the running time is spent in computing the simultaneous probabilities  
 383 of being below 0 in (7). This is done by calling the function `pmvnorm` of the R  
 384 package `mvtnorm` (Genz et al., 2019; Genz and Bretz, 2009). Uniform priors  
 385 are used for the parameters  $p_j$  and  $\beta_j$ , respectively on  $(0, 1)$  and  $(.25, 4)$ , while  
 386 PC priors are used for the parameters  $\mu_j$  and  $\alpha_j$ , as described in details in

**Algorithm 2** MCMC procedure

---

**Require:** Data; parent sequence; transform functions  $\varphi_j, j = 1, \dots, M$   
**Require:** Initial values and priors for all parameters  
**Require:** Number of iterations,  $N$

```

1: for  $i = 1$  to  $N$  do
2:   for each parameter  $\eta \in \{p, \mu, \beta, \alpha\}$  do
3:     for  $j = 1$  to  $M$  do
4:       Propose new  $\eta_j$  according to transition kernel
5:       Compute acceptance ratio,  $A$  using (14)
6:       Generate  $U \sim \mathcal{U}[0, 1]$ ; accept new  $\eta_j$  if  $(U \leq A)$ 
7:     end for
8:   end for
9:   for Borehole  $k = 1$  to  $n$  do
10:    Draw a move  $\in \{Split, Merge, Displace\}$  according to the probabilities
    ( $p_S, p_M, p_D$ )
11:    Check for feasibility within borehole  $k$ 
12:    if (move is feasible) then
13:      Draw uniformly one among all possible moves
14:      Compute acceptance ratio,  $A$  using (15)
15:      Generate  $U \sim \mathcal{U}[0, 1]$ ; accept the move if  $(U \leq A)$ 
16:    end if
17:  end for
18: end for

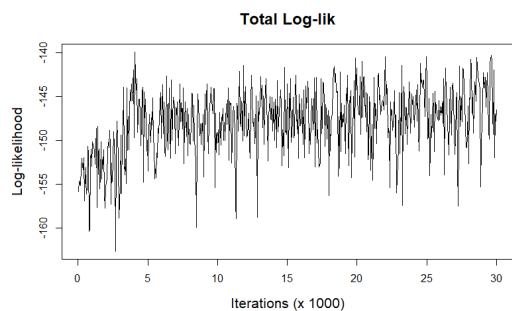
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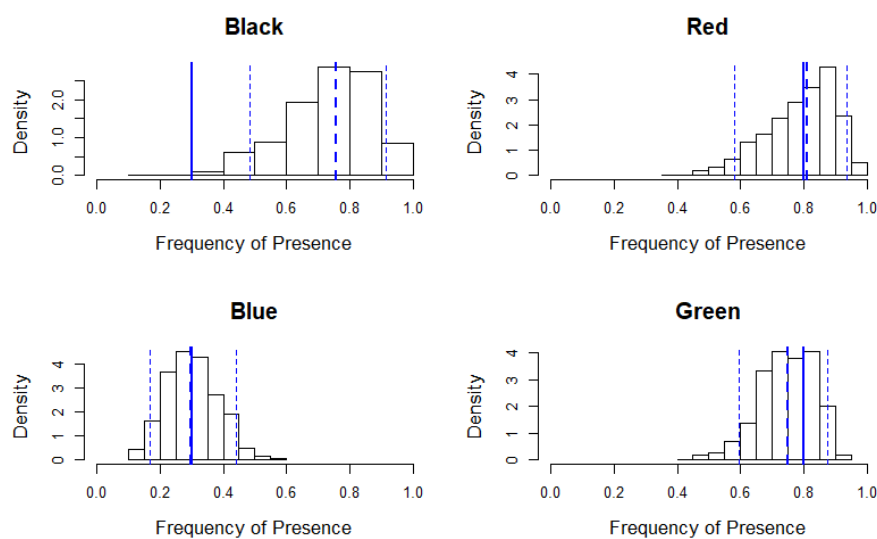
387 Sect. 4.2. Here, the setting was  $\epsilon_\alpha = \epsilon_\mu = 0.01$ , with  $\alpha_0 = 3$  and  $\mu_0 = 10$ . Algo-  
388 rithm 2 is run for 30,000 iterations, after a burn-in period of 2,500 iterations.  
389 Values of parameters are then sampled every 50 iterations. The proposals in  
390 the Metropolis-Hasting steps follow a uniform random walk with increments  
391 in  $[-0.4, 0.4]$  for  $\mu_j$  and  $\beta_j$ , in  $[-0.15, 0.15]$  for  $p_j$  and in  $[-3, 3]$  for the range  
392  $\alpha_j$ . With these choices, the observed acceptance ratio lies between 0.43 and  
393 0.57, depending on the parameters. This dataset being quite constrained, the  
394 acceptance ratio for exploring new configurations is only  $6.78 \cdot 10^{-5}$ .

## 395 5.1 Estimation of the Parameters

396 Figure 4 shows the complete log-likelihood as a function of the iterations. The  
397 mixing of the Markov chain is satisfactory and MCMC achieves convergence  
398 quite quickly. Figure 5 shows the posterior distribution of the frequency of  
399 each category. With the exception of the **Black** category, which was over-  
400 represented as already mentioned, the parameters  $p_j$  are very well estimated.  
401 Figure 6 shows the posterior cross-plot of the parameters  $\beta_j$  (resp.  $\alpha_j$ ) vs.  $\mu_j$ .  
402 One can see that there is some amount of negative correlation between  $\beta_j$  and  
403  $\mu_j$ , while there is some positive correlation between  $\alpha_j$  and  $\mu_j$ . These findings  
404 are quite consistent with the parametric form of the function  $\varphi(x) = \mu x^\beta$  on  
405 the one hand, and with the result obtained in Zhang (2004) regarding the  
406 simultaneous estimation of the range and variance of a Matérn random field  
407 on the other hand. One can observe that the posterior median is quite close to  
408 the true value and always within the 90% posterior credibility interval, at the  
409 exception of the range parameter for the **Black** category. For this category,

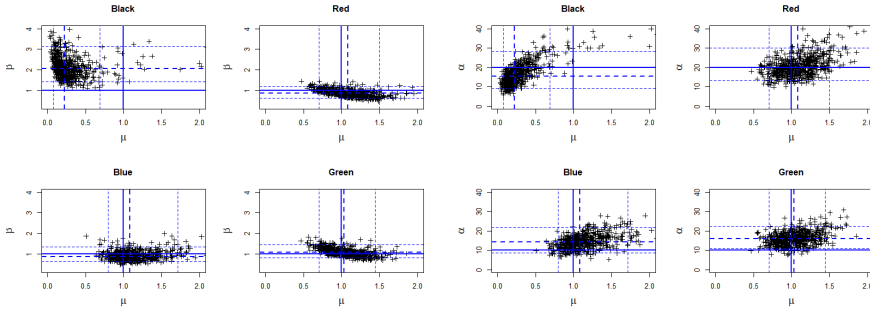


**Fig. 4** Complete log-likelihood as a function of iterations. The log-likelihood values are depicted every 50 iterations.



**Fig. 5** Posterior histograms of the frequencies  $p_j$ . Thick continuous line: true value of the parameter. Dashed thick line: posterior median. Dashed thin lines: posterior 0.05 and 0.95 quantiles.

410 it should be remembered that the observed frequency was over-represented  
 411 (0.58, as compared to 0.3) and that the average thickness was 0.31 as compared  
 412 to the theoretical expectation equal to 1.8. The maximum likelihood for the  
 413 parameters  $(p_j, \mu_j, \beta_j)$  is thus completely off the real values (0.3, 1, 1) as can  
 414 also be seen on Fig. 6, where  $\mu_j$  is under-estimated and  $\beta_j$  is over-estimated  
 415 (Fig. 6). Nonetheless, given the good performances in the other categories,  
 416 these results are quite promising considering that there are only 2 to 5 layers  
 417 per category and that there are only 12 synthetic boreholes.



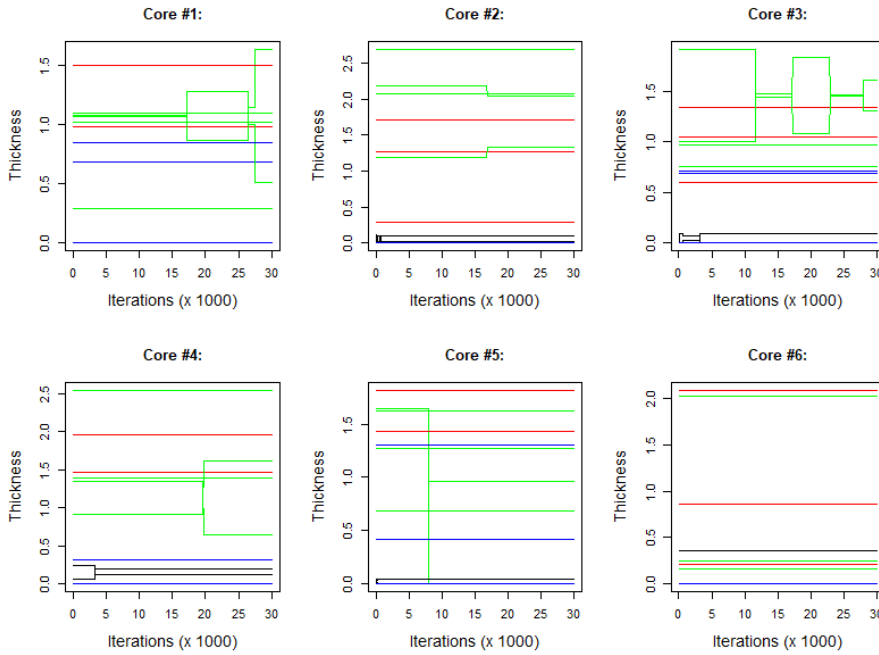
**Fig. 6** Left: posterior cross-plot of  $\beta_j$  vs.  $\mu_j$ . Right: posterior cross-plot of  $\alpha_j$  vs.  $\mu_j$ . Thick continuous line: true value of the parameter; Dashed thick line: posterior medians; dashed thin lines: posterior 0.05 and 0.95 quantiles.

## 418 5.2 Reconstruction of the Sequences

419 The observed sequence is not complete on most boreholes. Augmented se-  
 420 quences are created during the MCMC iterations. Since they can change along  
 421 the iterations, the MCMC algorithm allows us to explore different consistent  
 422 reconstructions. Figure 7 shows the thickness of the 15 layers as a function  
 423 of iterations for the first six synthetic boreholes. Each layer is color-coded ac-  
 424 cording to its category. Similar plots were obtained for the other boreholes,  
 425 but they are not shown here for the sake of concision. Firstly, it should be  
 426 noted that the thicknesses do not vary often and that the variability of the  
 427 thicknesses is quite different among the layers and among the boreholes. **Red**  
 428 layers show constant thickness because, in the parent sequence, **Red** layers are  
 429 separated by 4, respectively 6 layers (see Fig. 2). As a consequence, the condi-  
 430 tioning makes it impossible to *Merge* or *Split* any **Red** layers. The relative low  
 431 number of moves is due to the lateral correlations implied by the smoothness  
 432 parameter being equal to  $3/2$  and the range parameter being approximately  
 433 equal to  $1/3$  of the size of the domain. On boreholes #1 and #6, there is no  
 434 **Black** layer at all. The variations are not numerous and they concern mostly  
 435 the 6-layer sequence [**Green-Blue-Green-Blue-Green-Blue**] that allows some  
 436 exchanges of depth through successive moves. In particular, in boreholes #1  
 437 and #3 the actual sequence is [**Green-Blue-Green-Blue**], so that some of the  
 438 **Green** thickness can be exchanged between layers. Note that the total amount  
 439 of **Green** thickness remains always constant. On boreholes #2 to #5, some **Black**  
 440 layers are visible. The parent sequence is [**Black-Blue-Black**], but on borehole  
 441 #4 one of the observed thickness of **Blue** is 0. As a consequence, the observed  
 442 **Black** thickness can be shared between the two layers, or it can be attributed  
 443 to one layer only, the other being zero.

444 Figure 8 shows the thickness of layers #6 to #11 as a function of iterations  
 445 for each borehole intersecting the layer. It is the dual representation of Fig. 7.  
 446 Some layers have constant thickness across all boreholes, as it is the case for the  
 447 **Red** layer #7, which intersects 9 out of the 12 boreholes. On the three others,



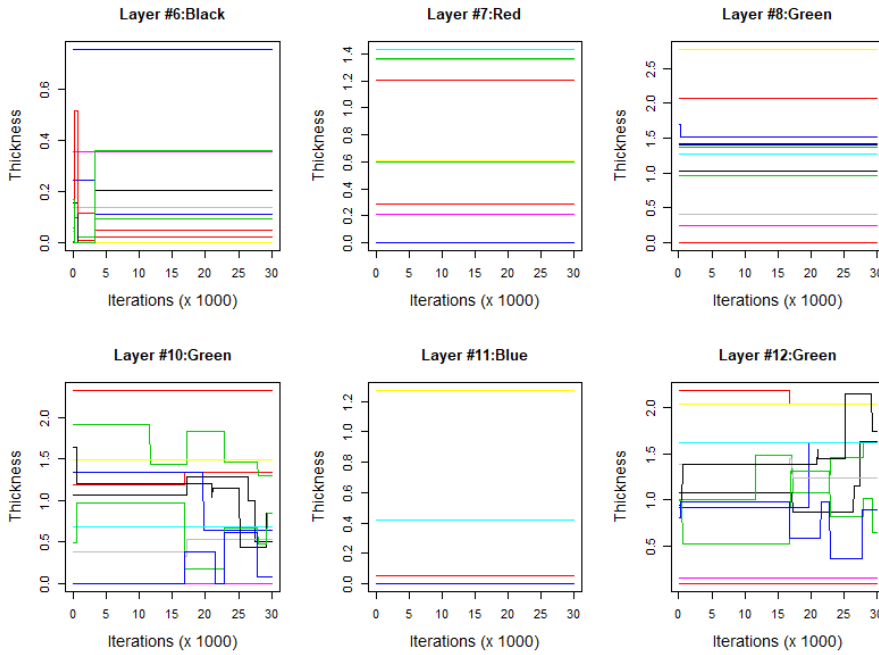


**Fig. 7** Thickness of different layers in synthetic boreholes #1 to #6 as a function of iterations. Layers are represented according to the color of the category they belong to.

448 the conditioning does not make it possible to *Merge* or *Split* the layer. In layers  
 449 #10 and #12, the situation is quite the opposite. Since the total thickness must  
 450 remain constant, variations on layers #10 and #12 are complementary for **Green**.  
 451 These layers are part of the [Green-Blue-Green-Blue] sequence from layer 10  
 452 to layer 13 already mentioned. This representation offers a complementary  
 453 view of the variations of this layer.

### 454 5.3 Conditional Simulations

455 Two ingredients are necessary in order to perform a simulation conditional on  
 456 the observed data. First, one needs all observed sequences to be coherently  
 457 completed in accordance with the parent sequence. Second, the simulation  
 458 requires parameters for  $\mu, \beta, p$  and  $\alpha$ . These must be jointly sampled from  
 459 the posterior distribution in a coherent way. Independent and identically dis-  
 460 tributed sets of augmented sequences and estimated parameters are accessible  
 461 by sampling from independently MCMC runs after the burn-in period. Alter-  
 462 natively, one can sample from the same MCMC run if the number of iterations  
 463 between two samples is large enough. The exact number depends on the mix-  
 464 ing properties of the MCMC algorithm. In practice, allowing a number of  
 465 iterations larger than the burn-in period is a safe enough option. The set of



**Fig. 8** Thickness of layers #6 to #12 as a function of iterations. Each borehole is represented with a different color.

466 parameters, together with the completed sequences corresponding to the high-  
 467 est likelihoods recorded, have been selected for conditional simulations. They  
 468 are depicted in Fig. 9. Both simulations honor perfectly the data at the bore-  
 469 holes (dashed vertical lines), but they show significantly different behaviors  
 470 away from the conditioning data.

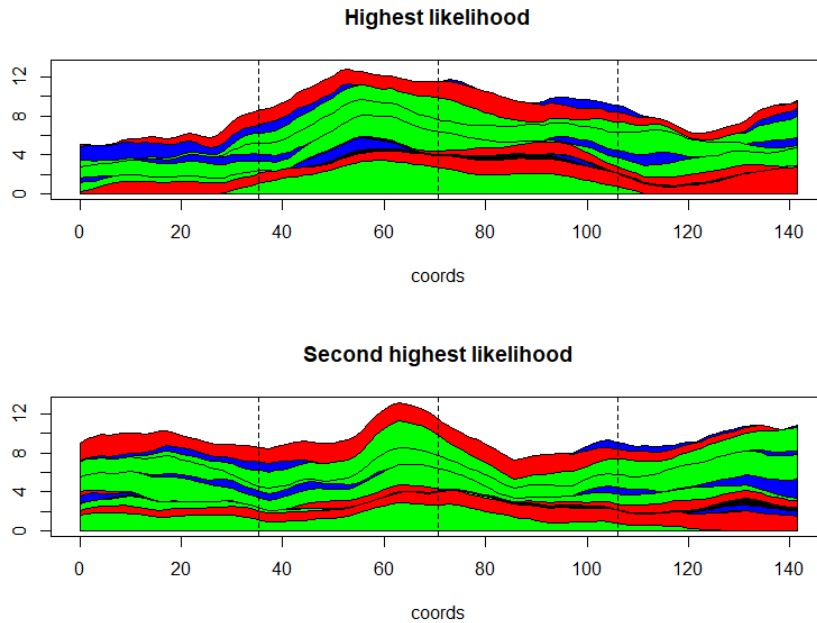
## 471 6 A Case Study: Deposition of Materials on an Aquifer

### 472 6.1 Study Area and Dataset Description

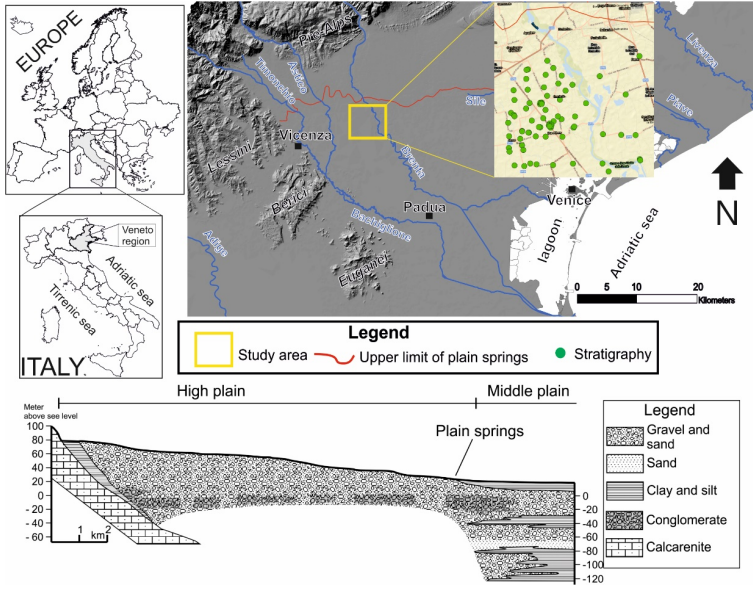
473 The study area (Fig. 10) is in the central part of the Venetian Plain (Italy),  
 474 on the Brenta megafan (principally on the right bank of the actual Brenta  
 475 River) of the Northern Padua district. In such an area, several rivers (Bac-  
 476 chiglione, Brenta, Astico and Timonchio) are responsible for the deposition of  
 477 a significant portion of the material, hundreds of meters thick, which forms  
 478 the subsoil of the Venetian Plain. Along the piedmont belt of the plain, fans  
 479 from adjacent rivers laterally penetrate gravelly alluvial fans. The result is  
 480 entirely gravelly subsoil throughout the thickness of the high Venetian Plain.  
 481 Because deeper fans often invade further areas of the high plain from the un-  
 482 differentiated gravel cover, the terminal parts of the fans extend downstream

483 for various distances, producing an alluvial cover that is no longer uniformly  
 484 gravelly, but is instead composed by alternating layers of gravel and silty clay  
 485 of swampy, lagoon or marine origin (Fabbri et al., 2016).

486 The data-set contains 24 boreholes drilled in a 5 km  $\times$  6 km region, with a  
 487 minimum distance between boreholes of 0.23 km (Fig. 11, top-left panel). Since  
 488 the maximum depth of the boreholes is highly variable, a depth window be-  
 489 tween the surface (from 35 m to 40 m above sea level) and 25 m above sea level  
 490 is selected. There are four categories **L**(imo) (Silt), **S**(abbia) (Sand), **G**(hiaia)  
 491 (Gravel), **A**(rgilla) (Clay) and the parent sequence, containing six layers, is:  
 492 [L-S-G-L-A-G]. Notice that, since there is only one layer for **S** and **A**, the associ-  
 493 ated thicknesses on the boreholes are known without ambiguity when present,  
 494 which is not necessarily the case for the thicknesses associated to **L** and **G**. From  
 495 two to four layers are observed on each borehole. One borehole contains an  
 496 observed sequence of length 4 and five boreholes contain an observed sequence  
 497 of length 3. The empirical estimates of the presence and the average thick-  
 498 nesses are shown in Table 3. The most observed categories are **S** followed by **L**  
 499 as measured by the proportion of presence (for **L**,  $\rho_j(0) = 22/(2 \times 24) = 0.46$ ).  
 500 The less observed category is **A**, with three records only.



**Fig. 9** Two conditional simulations. The completed sequences and the posterior parameters correspond to the most likely configuration of the MCMC run.



**Fig. 10** The study area of the real data example and the stratigraphy.

	L	S	G	A	Overall
Number of records	22	18	12	3	55
Proportion of presence, $p_j(0)$	0.46	0.75	0.25	0.13	0.38
Average thickness (in m), $\bar{T}_j$	0.73	2.25	3.89	1.10	1.94
Initial value, $\tau_j(0)$	0.10	-0.67	0.67	1.15	—
initial value, $\mu_j(0)$	0.96	2.06	6.52	2.21	—

**Table 3** Empirical estimates of presence, average thickness and initial values for  $\tau$  and  $\mu$ .

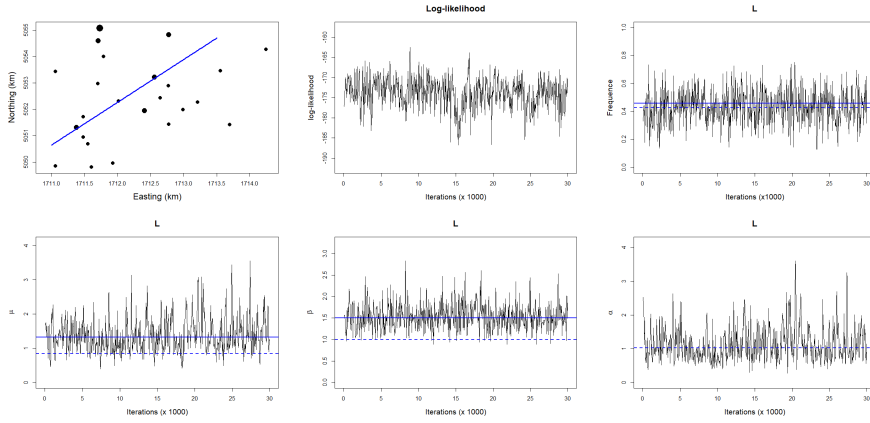
## 501 6.2 Model Setting

The empirical estimates are transformed into initial values for  $\tau_j$  and  $\mu_j$ , by setting initial values for  $\beta_j$  to  $\beta_j(0) = 1$ . Thus, for each category  $j$

$$\mu_j(0) = \frac{\bar{T}_j p_j(0)}{\phi(\tau_j(0))}, \quad \text{with } \tau_j(0) = \Phi^{-1}(1 - p_j(0)).$$

502 Preliminary tests (not reported here) showed that the likelihood computed  
 503 with a Matérn covariance function is almost always significantly larger with  
 504 a smoothness parameter  $\nu = 1/2$  than with  $\nu = 3/2$  or  $\nu = 5/2$ . Therefore,  
 505 the parameter  $\nu$  is set to  $1/2$ , corresponding to an exponential covariance  
 506 function, even though this covariance function corresponds to continuous but  
 507 non differentiable random surfaces. This point will be further discussed in Sect.  
 508 7. Initial values for the range are set to 1 km.

509 In this dataset, sequences are highly incomplete. As a consequence, the  
 510 MCMC algorithm needs to have good mixing properties in order to explore the  
 511 many possible augmented sequences that are compatible with the observations.



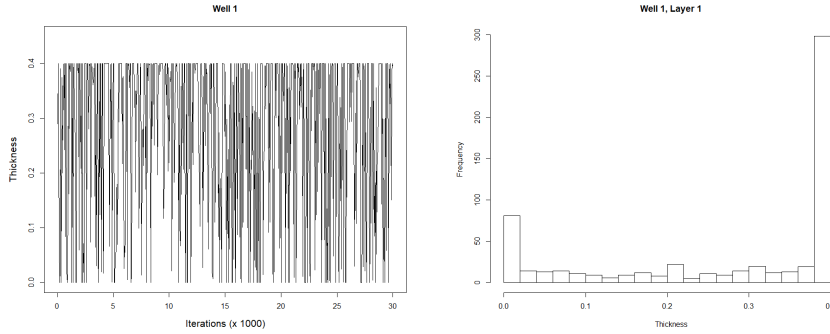
**Fig. 11** Location of the 24 boreholes analyzed in the Veneto dataset (top left); diameter is proportional to the number of thicknesses recorded (from 2 to 4); thick blue line: cross-section for conditional simulation. Then, from top to bottom and from left to right: total likelihood,  $p$ ,  $\mu$ ,  $\beta$  and  $\alpha$  as a function of iterations for category L. Continuous lines: posterior medians. Dashed lines: initial values.

512 Proposals follow a random walk with flat uninformative priors similar to that of  
 513 Sect. 4 for  $p_j$  and  $\beta_j$ . PC priors were used for  $\mu_j$  and  $\alpha_j$ , with  $\epsilon_\alpha = \epsilon_\mu = 0.01$   
 514 and  $(\alpha_0, \mu_0) = (.25, 10)$ . Algorithm 2 is run for 30,000 iterations, after a  
 515 burn-in period of 2,500 iterations. The values of the parameters are then  
 516 sampled every 50 iterations, so that  $m = 600$  posterior samples are collected.  
 517 The proposals in the Metropolis-Hasting steps follow a uniform random walk  
 518 with increments in  $[-0.4, 0.4]$  for  $\mu_j$  and  $\beta_j$ , in  $[-0.15, 0.15]$  for  $p_j$  and in  
 519  $[-0.2, 0.2]$  for the range  $\alpha_j$ . With these choices, the acceptance ratio for the  
 520 parameters was around 0.8. Although it is higher than recommended, it does  
 521 not appear to have a negative impact on the estimation procedure. Instead the  
 522 acceptance ratio of new thickness configurations was equal to 0.22 due to the  
 523 incompleteness of this data set. Figure 11 shows the values of the parameters  
 524  $p$ ,  $\mu$ ,  $\beta$  and  $\alpha$  as a function of iterations after burn-in, for category L. It  
 525 is quite clear that the chain is stationary with good mixing. Notice the difference  
 526 between the initial values and the posterior medians. Similar results have been  
 527 obtained for the other categories.

## 528 6.3 Results

### 529 6.3.1 Analysis of Thicknesses

530 When data belonging to the categories L and G are observed on the boreholes,  
 531 the recorded thickness might belong to a single layer or to two layers. For these  
 532 categories, the posterior thickness distribution might therefore look different  
 533 from the observed one. Figure 12 (left) shows how thicknesses of the first layer  
 534 L in borehole #1 vary along iterations thanks to the *Split*, *Merge* and *Displace*



**Fig. 12** Thickness of the first layer L in borehole # 1. Left: as a function of iterations. Right: posterior histogram.

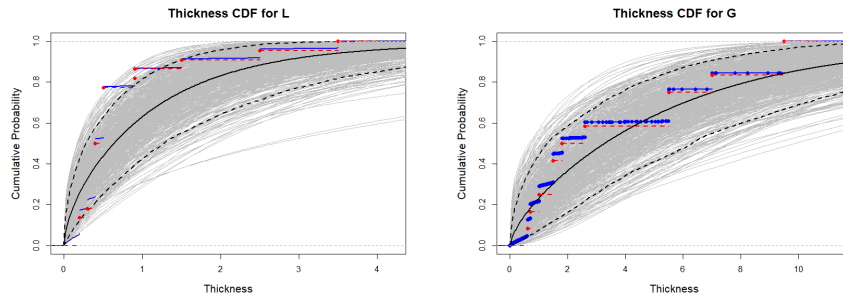
535 moves of the MCMC. On this borehole, the observed sequence is [L - A - G].  
 536 The measured thickness for L is equal to 0.4. Since the parent sequence is [L -  
 537 S - G - L - A - G] this thickness could correspond to the first layer only (case I),  
 538 to the fourth layer only (case II), or it could be shared between the two layers  
 539 (case III). Figure 12 (right) represents the posterior histogram of the thickness  
 540 in the first layer. Case I corresponds to 0.4, case II to 0 and case III to any  
 541 value in the interval  $(0, 0.4)$ . Frequencies computed along the iterations reveal  
 542 that case III is the most likely case, with an estimated probability of 0.47. The  
 543 probabilities of case I and case II are equal to 0.42 and 0.11, respectively. A  
 544 similar analysis can be performed easily on other boreholes and categories.

545 For a given category (for simplicity the index  $j$  is dropped), and for given  
 546 parameters  $(p, \mu, \beta)$ , the theoretical thickness cumulative distribution (TCD)  
 547 is

$$P(Z \leq z | p, \mu, \beta) = \int_{\tau}^{\tau + (z/\mu)^{1/\beta}} \frac{\phi(y)}{p} dy = \frac{\Phi(\tau + (z/\mu)^{1/\beta}) - \Phi(\tau)}{p}, \quad (16)$$

548 with  $\Phi(\tau) = 1 - p$ . The parameters are sampled every 50 iterations of the  
 549 MCMC, thereby mitigating the correlation between successive samples. At  
 550 each recorded iteration  $k = 1, \dots, m$ , the posterior samples  $p(k)$ ,  $\mu(k)$  and  
 551  $\beta(k)$  make it possible to compute a posterior theoretical TCD according to  
 552 (16). Those are represented in gray on Fig. 13 for categories L and G. The  
 553 ensemble of  $m$  posterior TCDs allows us to compute pointwise median and the  
 554 pointwise quantiles  $q_{0.05}$  and  $q_{0.95}$ , which are represented with black continuous  
 555 and dashed lines, respectively.

556 Empirical posterior TCD can alternatively be computed from the thickness  
 557 values recorded along the sampled iterations  $k = 1, \dots, m$ . In principle, em-  
 558 pirical and theoretical TCDs should match. Figure 13 shows the original and  
 559 posterior TCDs, respectively in red and blue. Thanks to the *Split*, *Merge* and  
 560 *Displace* movements, the posterior TCD is slightly smoother than the original  
 561 one since values intermediate to the observed ones are simulated.



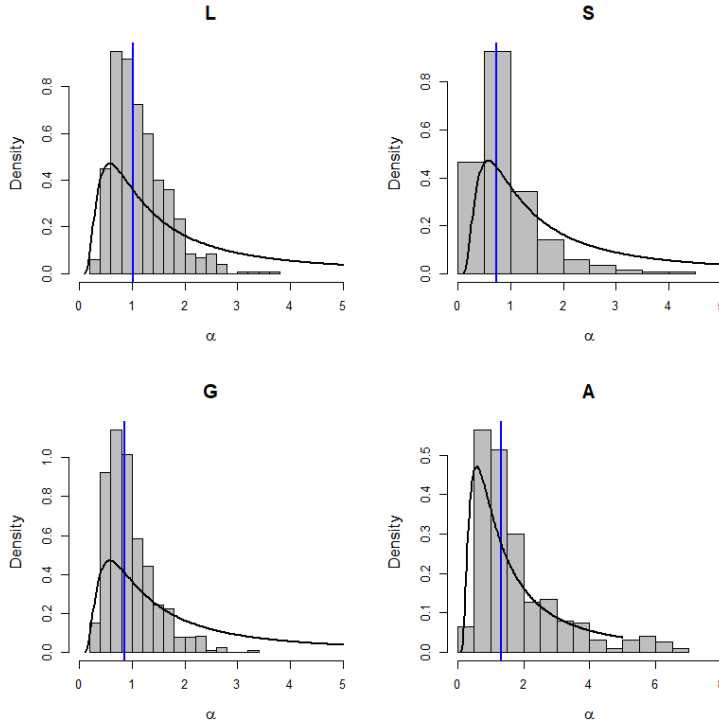
**Fig. 13** Thickness cumulative distributions (TCD). In gray: MCMC samples of the posterior theoretical TCD according to (16); Black continuous curve: pointwise posterior median TCD; Black dashed curves: pointwise posterior 0.05 and 0.95 posterior quantiles. Red dashed curve: TCD of the original data; Blue curve: TCD of the MCMC samples. Left: category L; Right: category G.

562 Overall, the match between the empirical and the theoretical TCD is very  
 563 satisfactory since the empirical curve is fully included in the envelope of the  
 564 MCMC samples for category G and is mostly included in the envelope for  
 565 category L.

### 566 6.3.2 Spatial Analysis and Conditional Simulation

567 Figure 14 shows the posterior histograms of the spatial range for the four  
 568 categories, with the prior density also shown. This figure indicates that the  
 569 prior has a heavy weight on the posterior distributions for each unit. However,  
 570 when a category is well informed (L and G), the posterior distribution is more  
 571 concentrated around the posterior median (indicated with a vertical blue line),  
 572 equal to 1.03, 0.73 and 0.85 for categories L, S and G, respectively. On the  
 573 contrary, category A has only three records. Since the likelihood contains very  
 574 little information, the posterior distribution is very close to the prior one. The  
 575 result of this analysis is that there is indeed a significant amount of spatial  
 576 correlations in the random fields modeling the thickness of the layers for all  
 577 categories but A.

578 Figure 15 shows two conditional simulations performed along the cross-  
 579 section depicted in Fig. 11 (top left). This cross-section has been chosen be-  
 580 cause it is close to three conditioning boreholes (shown with black vertical lines  
 581 on Fig. 15) with incomplete observed sequences that allow different thickness  
 582 configurations in category G. The color code is the following: red for L, blue  
 583 for S, green for G and black for A. The gray color corresponds to undefined  
 584 lithofacies below the last recorded layer. The first cross-section corresponds to  
 585 iteration 8,900 after burn-in, for which the likelihood was the highest along  
 586 the whole MCMC (log-likelihood is equal to  $-162.5$ ). Here, the G thickness is  
 587 entirely in layer # 6. The second cross-section corresponds to a configuration  
 588 where the G thickness is now shared between the two layers. Different shades  
 589 of green have been used to distinguish the two layers. This second configu-



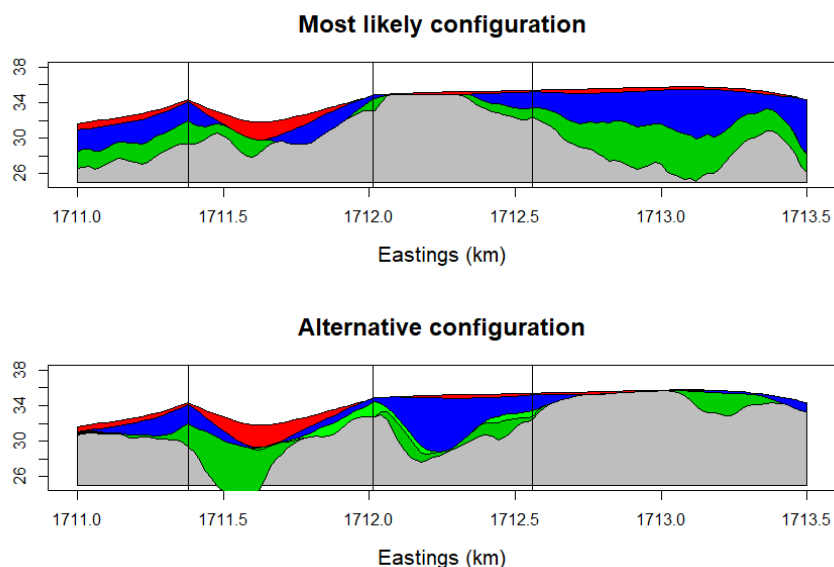
**Fig. 14** For each category, posterior histogram of the spatial range and prior distribution (continuous line). Blue vertical line: posterior median.

590 ration corresponds to the most likely configuration with shared thicknesses  
 591 between the two **G** layers (log-likelihood is equal to  $-171.3$ ). Notice that it is  
 592 significantly less likely than the first configuration, indicating that the data  
 593 is orders of magnitude less likely with the second configuration than with the  
 594 first one. Notice also that the cross-sections are quite different when moving  
 595 away from the conditioning boreholes. The parameters corresponding to these  
 596 two configurations are reported in Table 4.

	First configuration				Second configuration			
	Log-likelihood = $-162.5$				Log-likelihood = $-171.3$			
	L	S	G	A	L	S	G	A
$p_j$	0.40	0.81	0.23	0.05	0.45	0.64	0.46	0.48
$\mu_j$	1.29	1.80	6.99	2.20	1.98	2.02	5.01	11.03
$\beta_j$	1.54	1.59	1.01	0.50	1.42	1.41	1.39	1.76
$\alpha_j$	1.25	0.29	0.78	0.71	2.03	0.54	0.43	3.58

**Table 4** Parameters corresponding to the two configurations shown in Fig. 15.





**Fig. 15** Two cross-sections along the line shown in Fig. 11 (top left). Notice that there are two different layers for  $G$  in the bottom cross-section.

## 597 7 Concluding Remarks

598 In this paper a new rule-based approach for simulating depositional sequences  
 599 of surfaces conditionally to lithofacies thickness data has been presented. A dis-  
 600 tinctive feature of this approach is that it takes into account in a coherent way  
 601 the different amount of information along horizontal and vertical dimensions  
 602 that are usually contained in borehole datasets: few cores and, consequently,  
 603 few horizontal information but complete information along the depth.

604 This is achieved by supposing that there exists a common lithological se-  
 605 quence of facies that is compatible with the observed data. Moreover the se-  
 606 quence is supposed to be known in advance. The facies thickness, which is  
 607 non-negative, is modeled by means of a truncated and transformed stationary  
 608 Gaussian field. In principle, other non-negative random fields could be con-  
 609 sidered, but this choice made it possible to exploit the flexibility of Gaussian  
 610 random fields in the selection of the covariance functions with different degrees  
 611 of smoothness. The evaluation of the likelihood is made possible thanks to the  
 612 Gaussian framework for which well known methods and efficient computing  
 613 tools are available.

614 A data augmentation algorithm, coupled with a MCMC algorithm, is em-  
 615 ployed for learning the parameters of the stochastic model from borehole data.  
 616 A very interesting feature of the proposed algorithm is that the exploration  
 617 of all different configurations that are compatible with the available data is  
 618 possible. Thanks to the MCMC approach and the Bayesian framework, it as-

619 sociates a likelihood to each of the possible realizations corresponding to a set  
620 of parameters. From those, as shown in Sect. 6.3, one can assess an empirical  
621 probability for each different configuration, select the most likely configura-  
622 tions and compute many other statistics of interest to the user. The algorithm  
623 requires multiple (to the order of  $M \times n$ ) evaluations of the joint probability  
624 of a Gaussian vector being below a given threshold. The current implementa-  
625 tion in R uses the `mvtnorm` package (Genz et al., 2019) that handles vectors  
626 with a few dozens of coordinates rather easily. It starts to slow down quite  
627 significantly around 100 coordinates and is unable to cope with more than  
628 1,000 coordinates. Further research is thus required if the number of boreholes  
629 goes from moderate to high or very high. One possible choice could be the  
630 approximation proposed in Martinetti and Geniaux (2017), but the impact of  
631 using a less precise approximation remains to be evaluated.

632 A too small dataset entails difficulties in specifying the regularity and the  
633 range of the covariance function, as was shown with category A that has only  
634 three records. It was found in the present work that parameters were reason-  
635 ably well estimated with 15 records per category. On the other hand, as the  
636 data set gets larger and denser (e.g. when the horizontal distance between  
637 nearest neighbor boreholes becomes a small fraction of the range parameter)  
638 the likelihood will get more peaked around local maxima, thereby decreasing  
639 the mixing of the MCMC. In this case, exploring all configurations coherent  
640 with the parent sequence is likely to become more difficult. Longer chains and  
641 multiple chains starting from very different initial configurations will probably  
642 be necessary.

643 Several assumptions and restrictions have been made in this work, which  
644 can be lifted in order to generalize this work. The stationarity assumption,  
645 which has proved appropriate here, could be relaxed and the parameters could  
646 be easily modified to take covariates into account. Only a few half-integer  
647 values of the smoothness parameters have been considered, and the fitting  
648 of this parameter was done outside the MCMC machinery. In principle, the  
649 smoothness parameter could be different for different facies and it could be  
650 estimated in the Bayesian framework, just as any other parameter. Estimating  
651 simultaneously the three parameters of the Matérn covariance in a Bayesian  
652 context is known to be extremely difficult. When there are only few data, this  
653 was made possible thanks to the PC priors (Fuglstad et al., 2019). Currently, to  
654 the best of our knowledge, the simultaneous PC prior for  $(\nu, \alpha, \sigma^2)$  for Matérn  
655 covariance is unknown. Finding such PC priors is left for further research.

656 Currently, independent MCMCs are launched, one for every possible value  
657  $\nu \in \{1/2, 3/2, 5/2\}$ . The one with the highest likelihood and the best mixing  
658 is selected and  $\nu$  is fixed at that value. When analyzing the data from the  
659 Venetian plain, it was found that  $\nu = 1/2$  was best, despite the fact that  
660 the associated thicknesses (and thus surfaces) are mean-square continuous but  
661 not differentiable. One could have imposed  $\nu = 3/2$ , but at the cost of a very  
662 short spatial range implying almost no spatial correlation. Whether one should  
663 let the data speak or impose a model for the regularity is a debate. Here, a  
664 data-driven approach was chosen.

665 Finally, the function that transforms the Gaussian values to thicknesses  
666 was chosen to be a power function, but any other positive function could be  
667 used.

668 One information that is often available in real applications and on much  
669 more points than boreholes is the nature of the facies on the surface. It is  
670 possible to incorporate such information at the cost of small changes in the  
671 method. At a given location  $s$  where this information is available, one could  
672 consider that the facies of the upper layer, say facies  $j$ , is known and has a  
673 positive thickness. The conditioning data would therefore be that  $W_{\text{upper}}(x) >$   
674  $\tau_j$ . This conditioning can easily be handled within our MCMC procedure. At  
675 this location, there would be no conditioning for the other layers.

676 The proposed approach depends on the existence and the knowledge of a  
677 common lithological sequence of facies that is compatible with the observed  
678 data. If the sequence is unknown, it is possible to derive it from the data,  
679 possibly by imposing some restriction, such as minimum length. This problem  
680 has not been tackled here, since it has been considered beyond the scope of  
681 this work. However it is worth mentioning that the approach presented here  
682 can be modified to account for several different parent sequences with their  
683 associated prior probabilities.

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688 **A A longer example of incomplete sequence**

**Table 5** A longer and more complex example of a parent sequence  $C=[\text{Blue-Red, Blue-Green-Blue-Red-Green-Blue}]$  with respect to a recorded sequence  $C^o$  and  $T^o$ . Only nine compatible augmented sequences are reported.

Parent		Recorded		Compatible augmented sequences								
C	C <sup>o</sup>	T <sup>o</sup>	C <sup>a</sup>	T <sup>a</sup>	Z <sup>a</sup>	C <sup>a</sup>	T <sup>a</sup>	Z <sup>a</sup>	C <sup>a</sup>	T <sup>a</sup>	Z <sup>a</sup>	
Blue	Blue	T <sub>1</sub> <sup>o</sup>	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	
Red	Red	T <sub>2</sub> <sup>o</sup>	Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>		T <sub>1</sub> <sup>o</sup>	0	
Blue	Green	T <sub>3</sub> <sup>o</sup>		T <sub>2</sub> <sup>o</sup>	0		T <sub>2</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0	
Green	Blue	T <sub>4</sub> <sup>o</sup>	Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>		T <sub>2</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0	
Blue	-	-	Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>		T <sub>2</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0	
Red	-	-		T <sub>4</sub> <sup>o</sup>	0		T <sub>2</sub> <sup>o</sup>	0	Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	
Green	-	-		T <sub>4</sub> <sup>o</sup>	0	Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>	Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>	
Blue	-	-		T <sub>4</sub> <sup>o</sup>	0	Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>	Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>	
			Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	
			Red	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>		T <sub>1</sub> <sup>o</sup>	0	
				T <sub>1</sub> <sup>o</sup>	0		T <sub>2</sub> <sup>o</sup>	0	Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	
				T <sub>1</sub> <sup>o</sup>	0	Green	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>	Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>	
				T <sub>1</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>	
			Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>		T <sub>1</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0	
			Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>	Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>		T <sub>1</sub> <sup>o</sup>	0	
			Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>	Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>	Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	
			Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup>	
				T <sub>1</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0	
				T <sub>1</sub> <sup>o</sup>	0	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	
				T <sub>1</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0		T <sub>1</sub> <sup>o</sup>	0	
			Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>		T <sub>1</sub> <sup>o</sup>	0	Blue	T <sub>1</sub> <sup>o</sup>	T <sub>1</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	
			Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	Red	T <sub>2</sub> <sup>o</sup>	T <sub>2</sub> <sup>o</sup> - T <sub>1</sub> <sup>o</sup>	
			Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>	Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>	Green	T <sub>3</sub> <sup>o</sup>	T <sub>3</sub> <sup>o</sup> - T <sub>2</sub> <sup>o</sup>	
			Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>	Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>	Blue	T <sub>4</sub> <sup>o</sup>	T <sub>4</sub> <sup>o</sup> - T <sub>3</sub> <sup>o</sup>	

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